Ocean Mixed Layers



Fig. 2. Profile taken during convective equilibrium on PATCHEX; later in the same night as shown in Fig. 1. Panel a. Profiles of ε , θ , salinity and σ_{θ} . The shading is ε , estimated in 0.5 m bins: θ and σ_{θ} have been processed with a 0.8 m triangular filter. Panel b. Thorpe displacements; the instability extending to 0.48 MPa shows active convection reaching the seasonal thermocline.

0.5 Mpa ≈ 50m depth





- Properties close to the surface of the ocean tend to be well mixed and relatively uniform with depth.
- Depth of mixed layer depends on surface forcing (wind and buoyancy) and entrainment from below and can vary from O(1 meter) to in excess of 1000m.
- Strong warming or freshwater input leads to a shallow layer
- Strong cooling leads to a deep layer
- Properties exchanged with the atmosphere (heat, freshwater, gases) are quickly mixed throughout the layer of active mixing (a few minutes to a few hours)
- Entrainment at base of mixed layer brings nutrients into the surface layer.
- Light received by phytoplankton depends on the depth of the layer and level of turbulence.

Mixed Layer Depth, h

Definition based on change from surface:

 $\Delta \sigma(z) = \sigma(z) - \sigma(z=0)$

 h_{σ} depth when $\Delta\sigma$ crosses some threshold (say 0.03kgm-3; needs to be refined depending on situation)

Can define in terms of temperature

 $\Delta T=T(z=0)-T(z)$

 h_T depth when ΔT crosses some threshold (say 0.2°C)

NOTE: h_{σ} not always the same as h_{τ}



Barrier layer: Lindstrom et al 1987





Fig. 5. Diagram showing depth zones in a typical diurnal mixed layer cycle.

Basics: warming







PE 🖡



Requires TKE to mix heat down. PE 1 Change in SST dependent on mixing within the layer and entrainment from below.

Basics: cooling







Indian Ocean, 8°S October 2011



Temperature

Shear

Turbulence

Indian Ocean, 8°S October 2011



Salinity



Shear

Turbulence





Figure 1 Seasonal evolution of temperature in the midlatitude upper ocean. Shallow, warm mixed layers during spring-summer alternate with cold, deep ones during fall-winter. (Kantha and Clayson, 2000.)

mld [m], (file name = mld_DR003_c1m_reg2.0.nc), mld ≈ Isopycnal Layer Depth





90 135

45

-45 0

130 -135

-- 90



45 0

-135 `-90

45

90 135





Additional influences: Langmuir Circulation





Additional influences: Mixed Layer Instability



FIG. 2. Temperature (°C) during two typical simulations of a ML front spinning down: (a)–(c) no diurnal cycle and (d)–(f) with diurnal cycle and convective adjustment. (Black contour interval = 0.01°C; white contour interval = 0.1°C)

Tendency to re-stratify the upper ocean: impact on biological production Sub-mesoscale structures in a regional implementation of HYCOM forced with a regional atmospheric model: Jan 15, 2010





ΔSST

Mean N top 30m

Connection with the interior: Subduction and Ventilation



Subtropical gyres: The combination of the equatorward flow plus poleward increase in mixed layer depth leads to the subduction of water



Figure 2 A schematic diagram illustrating the seasonal cycle of the mixed layer following the movement of a water column. The mixed layer thins in spring and summer, and thickens again in autumn and winter. If there is an overall buoyancy input, the end of winter mixed layer becomes lighter and thinner from one year to the next (as depicted here). Consequently, fluid is subducted irreversibly from the mixed layer into the main or permanent thermocline. The mixed layer thickness is marked by the thick dashed line, isopycnals ρ_m outcropping at the end of winter into the mixed layer by the full lines, and the isopycnal identifying the base of the seasonal thermocline by the short dashed line. The annual subduction rate, S_{ann} , determines the vertical spacing between the isopycnals subducted from the mixed layer in March for consecutive years 1 and 2 (τ_{year} is one year).



Propagation of cold/fresh anomalies on σ_{θ} 25-25.5 kg m⁻³ (Data: Argo. Orange dashed line: Feb outcrop of surface)

Sasaki et al GRL 2010