# **Chapter 5**

# **Physical and** thermodynamic data

# **Table of Contents**



#### **1. Values for various fundamental constants**

- $R = 8.314472(15)$  J K<sup>-1</sup> mol<sup>-1</sup>,
- $F = 96485.3399(24)$  C mol<sup>-1</sup>,
- $0^{\circ}$ C = 273.15 K (defined),
- 1 atm = 101325 Pa (defined)

(from http://physics.nist.gov/cuu/Constants/index.html). For each constant, the standard uncertainty in the least significant digits is given in parentheses.

# **2. Atomic weights**

Atomic weights are quoted here (Table 1) to 5 significant figures unless the dependable accuracy is more limited by either the combined uncertainties of the best published atomic weight determinations, or by the variability of isotopic composition in normal terrestrial occurrences (elements **annotated r**). The last significant figure of each tabulated value is considered reliable to at least  $\pm 1$ except when a larger single-digit uncertainty is inserted in parentheses following the atomic weight. In many cases the "official" values (IUPAC, 2006) have more than five significant figures and the precision is considerably better than  $\pm 1$ at the precision shown here. In a few cases the truncation to five significant figures will mean that the true confidence interval is strongly asymmetric relative to the values shown here. Neither the highest nor the lowest actual atomic weight of any normal sample is thought likely to differ from the tabulated value by more than the assigned uncertainty. However, the tabulated values do not apply either to samples of highly exceptional isotopic composition arising from most geological occurrences (elements **annotated g**) or to those whose isotopic composition has been artificially altered. Such might even be found in commerce without disclosure of that modification (elements **annotated m**). Elements that have no stable isotope are excluded from this table except for Th, Pa, and U which do have a characteristic terrestrial isotope composition.



**Table 1** IUPAC (2006) table of standard atomic weights abridged to five significant figures. Scaled to the relative atomic mass  $A(^{12}C) = 12$  (includes 2007 revisions).

 $\overline{\phantom{0}}$ 



#### **Table 1** Continued.





\* These elements have no stable isotope.

# **3. Vapor pressure of water**

#### **3.1 Pure water**

The vapor pressure of pure water  $(p_{\sigma})$  over the temperature range 273 to 647 K (*i.e.*, from the triple point to the critical point of water) is represented by the polynomial (Wagner and Pruß, 2002)

$$
\ln\left(\frac{p_{\sigma}}{p_{c}}\right) = \frac{T_{c}}{T}\left(a_{1}\theta + a_{2}\theta^{1.5} + a_{3}\theta^{3} + a_{4}\theta^{3.5} + a_{5}\theta^{4} + a_{6}\theta^{7.5}\right)
$$
 (1)

where  $\theta = (1 - T/T_c)$ ,  $T_c = 647.096$  K and  $p_c = 22.064$  MPa. The coefficients are:



At 25°C (298.15 K),  $p_{\sigma}$  = 3.1698 kPa.

#### **3.2 Sea water**

For sea water, the vapor pressure is related to that of pure water,

$$
p_{\sigma} \left( \text{s/w} \right) \approx p_{\sigma} \exp \left( -0.018 \phi \sum_{\text{B}} m_{\text{B}} / m^{\circ} \right) \tag{2}
$$

where

 $p_{\sigma}$  is the vapor pressure of pure water (1),

 $\sum_{p} m_{\text{B}}$  is the total molality of dissolved species (3),

 $\phi$  is the osmotic coefficient of sea water (4),  $m^{\circ}$  = 1 mol kg-H<sub>2</sub>O<sup>-1</sup>.

For sea water (of the composition detailed in section 6.1),

$$
\sum_{B} m_{B}/m^{\circ} = \frac{31.998S}{10^3 - 1.005S}
$$
 (3)

and the osmotic coefficient at 25°C is given by (Millero, 1974)

$$
\phi = 0.90799 - 0.08992 \left( \frac{\frac{1}{2} \sum_{B} m_{B}}{m^{\circ}} \right) + 0.18458 \left( \frac{\frac{1}{2} \sum_{B} m_{B}}{m^{\circ}} \right)^{2} - 0.07395 \left( \frac{\frac{1}{2} \sum_{B} m_{B}}{m^{\circ}} \right)^{3} - 0.00221 \left( \frac{\frac{1}{2} \sum_{B} m_{B}}{m^{\circ}} \right)^{4} \tag{4}
$$

for  $I = 0.3 - 0.8$  mol kg-H<sub>2</sub>O<sup>-1</sup> (*S* = 16–40). (Although strictly speaking,  $\phi$  is a function of temperature as well as of  $I$ , the change between 0 and  $40^{\circ}$ C is typically about 1% and can be ignored here.)

At 25<sup>o</sup>C and *S* = 35,  $p_{\sigma}$  (s/w) = 3.1106 kPa.

#### **4. Densities of various substances**

#### **4.1 Air-saturated water**

The density of air-saturated water in the temperature range 5 to 40°C—*i.e.*, avoiding the temperature of maximum density—is given by the expression (Jones and Harris, 1992)

$$
\rho_{\rm w}/(\text{kg m}^{-3}) = 999.84847 + 6.337563 \times 10^{-2} (t/^{\circ}\text{C}) - 8.523829 \times 10^{-3} (t/^{\circ}\text{C})^2 + 6.943248 \times 10^{-5} (t/^{\circ}\text{C})^3 - 3.821216 \times 10^{-7} (t/^{\circ}\text{C})^4.
$$
 (5)

At 25°C,  $\rho_{\text{W}}$  = 997.041 kg m<sup>-3</sup> = 0.997041 g cm<sup>-3</sup>.

#### **4.2 Sea water**

The density of sea water in the temperature range 0 to  $40^{\circ}$ C (IPTS  $68^{\circ}$ ) and the salinity range 0 to 42 is given by the expression (Millero and Poisson, 1981)

$$
\rho_{SW}/(\text{kg m}^{-3}) = \rho_{SMOW}/(\text{kg m}^{-3}) + AS + BS^{1.5} + CS^{2}
$$
 (6)

where $2$ 

<u>.</u>

$$
\rho_{\text{SMOW}} / (\text{kg m}^{-3}) = 999.842594 + 6.793952 \times 10^{-2} (t/^{\circ}\text{C}) \n- 9.095290 \times 10^{-3} (t/^{\circ}\text{C})^2 + 1.001685 \times 10^{-4} (t/^{\circ}\text{C})^3 \n- 1.120083 \times 10^{-6} (t/^{\circ}\text{C})^4 + 6.536332 \times 10^{-9} (t/^{\circ}\text{C})^5 , \quad (7)
$$

$$
A = 8.24493 \times 10^{-1} - 4.0899 \times 10^{-3} (t/^{\circ}\text{C})
$$
  
+7.6438 \times 10^{-5} (t/^{\circ}\text{C})^2 - 8.2467 \times 10^{-7} (t/^{\circ}\text{C})^3 + 5.3875 \times 10^{-9} (t/^{\circ}\text{C})^4, (8)

$$
B = -5.72466 \times 10^{-3} + 1.0227 \times 10^{-4} (t/^{\circ}\text{C}) - 1.6546 \times 10^{-6} (t/^{\circ}\text{C})^{2},\tag{9}
$$

$$
C = 4.8314 \times 10^{-4},\tag{10}
$$

*t* is on IPTS 68 and *S* is the salinity.

At 25°C (IPTS 68) and  $S = 35$ ,  $\rho_{SW} = 1023.343$  kg m<sup>-3</sup>.

#### **4.3 Sodium chloride solutions**

The density of a solution of sodium chloride in water in the temperature range 0 to 50°C (see Footnote 1) and the concentration range 0.1 mol kg- $H_2O^{-1}$  to saturation can be calculated from the expression (Lo Surdo *et al.*, 1982)

$$
\frac{10^{3}(\rho(\text{NaCl}) - \rho_{\text{SMOW}})}{g \text{ cm}^{-3}} = \left(\frac{m}{m^{\circ}}\right) \left[45.5655 - 0.2341\left(\frac{t}{\text{°C}}\right)\n+3.4128 \times 10^{-3}\left(\frac{t}{\text{°C}}\right)^{2} - 2.7030 \times 10^{-5}\left(\frac{t}{\text{°C}}\right)^{3} + 1.4037 \times 10^{-7}\left(\frac{t}{\text{°C}}\right)^{4}\right]\n+\left(\frac{m}{m^{\circ}}\right)^{3/2} \left[-1.8527 + 5.3956 \times 10^{-2}\left(\frac{t}{\text{°C}}\right) - 6.2635 \times 10^{-4}\left(\frac{t}{\text{°C}}\right)^{2}\right]\n+\left(\frac{m}{m^{\circ}}\right)^{2} \left[-1.6368 - 9.5653 \times 10^{-4}\left(\frac{t}{\text{°C}}\right) + 5.2829 \times 10^{-5}\left(\frac{t}{\text{°C}}\right)^{2}\right]\n+0.2274\left(\frac{m}{m^{\circ}}\right)^{5/2}
$$
\n(11)

$$
_{90}/^{\circ}C = 0.0002 + 0.99975 t_{68}/^{\circ}C
$$

<sup>1</sup> The International Practical Temperature Scale of 1968 (IPTS 68) has been superseded by the International Temperature Scale of 1990 (ITS 90). A simple equation can be used to relate the two over the oceanographic temperature range 0 to 40°C (Jones and Harris, 1992):

 $t_{90}/^{\circ}C = 0.0002 + 0.99975 t_{68}/^{\circ}C$ .<br><sup>2</sup> SMOW—Standard Mean Ocean Water (Craig, 1961)—is pure water with a specified isotopic composition and free of dissolved gases.

where  $t$  is the temperature (IPTS  $68$ ),  $m$  the molality of the sodium chloride solution, and  $m^{\circ} = 1$  mol kg-H<sub>2</sub>O<sup>-1</sup>.  $\rho_{SMOW}$  is calculated using equation (7)—see Footnote 3.

To convert the concentration of sodium chloride expressed in mol kg-soln<sup>-1</sup> $C(NaCl)$ —to mol kg-H<sub>2</sub>O<sup>-1</sup>, use the expression

$$
\frac{m(NaCl)}{m^{\circ}} = \frac{10^3 C(NaCl)}{10^3 - 58.443 C(NaCl)}.
$$
(12)

At 25°C and  $m = 0.725$  mol kg-H<sub>2</sub>O<sup>-1</sup>,  $\rho$  (NaCl) = 1.02581 g cm<sup>-3</sup>.

#### **4.4 A mixture of NaCl and HCl at 25°C**

An expression based on Young's rule has been developed for the density of a mixture of NaCl and HCl at a temperature of  $25^{\circ}C^4$  (Millero, personal communication):

$$
\rho_{\text{mix}}(25 \text{ °C}) = \frac{\rho_{\text{w}}(25 \text{ °C}) \left[10^3 + m_{\text{T}} \left(m(\text{HCl}) + m(\text{NaCl})\right)\right]}{10^3 + \varphi_{\text{mix}} \left(m(\text{HCl}) + m(\text{NaCl})\right) \left(\rho_{\text{w}}(25 \text{ °C})\right)}
$$
(13)

where  $\rho_{\rm W}(25^{\circ}\rm C) = 0.99704 \text{ g cm}^{-3}$ ,

$$
m_{\rm T} = \frac{36.46m(\text{HCl}) + 58.44m(\text{NaCl})}{m(\text{HCl}) + m(\text{NaCl})},\tag{14}
$$

$$
\varphi_{\text{mix}} = \frac{m(\text{HCI})\varphi_{\text{HCI}} + m(\text{NaCl})\varphi_{\text{NaCl}}}{m(\text{HCI}) + m(\text{NaCl})},\tag{15}
$$

and the apparent molar volumes at 25°C,

$$
\varphi_{\text{HCl}} = 17.854 + 1.460\sqrt{m - 0.307m} \,,\tag{16}
$$

$$
\varphi_{\text{NaCl}} = 16.613 + 1.811\sqrt{m} + 0.094m\tag{17}
$$

where  $m = m(HCl) + m(NaCl)$ . The concentrations  $m(HCl)$  and  $m(NaCl)$  are in mol kg-H<sub>2</sub>O<sup>-1</sup>.

At  $m(HCl) = 0.2$  mol kg-H<sub>2</sub>O<sup>-1</sup> and  $m(NaCl) = 0.5$  mol kg-H<sub>2</sub>O<sup>-1</sup>,  $\rho_{mix}(25^{\circ}C) =$  $1.02056$  g cm<sup>-3</sup>.

#### **4.5 Various solids**

From data of Weast (1975),

$$
Potassium chloride: \qquad \rho (KCI) = 1.984 \text{ g cm}^{-3} \tag{18}
$$

 $\overline{a}$ 

<sup>3</sup> In the original paper of Lo Surdo *et al.* (1982) an alternate formulation based on the work of Kell (1975) was used for the density of pure water. The difference is  $\leq 10^{-5}$ over the temperature range 0–50°C.

At temperatures other than 25°C, equation (13) is not reliable.

- Sodium chloride:  $\rho(NaCl) = 2.165 \text{ g cm}^{-3}$  (19)
- Sodium carbonate:  $\rho(Na_2CO_3) = 2.532 \text{ g cm}^{-3}$  (20)

Sodium sulfate:  $\rho(Na_2SO_4) = 2.68 \text{ g cm}^{-3}$  (21)

Sodium tetraborate decahydrate (borax):

$$
\rho \left( \text{Na}_2\text{B}_4\text{O}_7 \cdot 10\text{H}_2\text{O} \right) = 1.73 \text{ g cm}^{-3} \tag{22}
$$

2-amino-2-hydroxymethyl-1,3-propanediol ("TRIS"):

$$
\rho(H_2NC(CH_2OH)_3) = 1.35 \text{ g cm}^{-3}
$$
 (23)

2-aminopyridine: 
$$
\rho(C_5H_6N_2) = 1.24 \text{ g cm}^{-3}
$$
 (24)

#### **5. Virial coefficients for carbon dioxide**

#### **5.1 Pure carbon dioxide gas**

The first virial coefficient of  $CO<sub>2</sub>$  is given by the expression (Weiss, 1974)

$$
\frac{B(CO_2, T)}{cm^3 mol^{-1}} = -1636.75 + 12.0408 \left(\frac{T}{K}\right) - 3.27957 \times 10^{-2} \left(\frac{T}{K}\right)^2 + 3.16528 \times 10^{-5} \left(\frac{T}{K}\right)^3
$$
\n(25)

where 265 < *T*/K < 320.

At 25<sup>o</sup>C (298.15 K),  $B(CO_2, 298.15 K) = -123.2 \text{ cm}^3 \text{ mol}^{-1}$ .

#### **5.2 Carbon dioxide in air**

For a binary mixture (B–C), it is possible to define a cross-virial coefficient:

$$
\delta_{\rm B-C} = B_{\rm BC} - \frac{1}{2} (B_{\rm BB} - B_{\rm CC}) \,. \tag{26}
$$

Weiss (1974) used the Lennard–Jones (6–12) potential to calculate values for the cross-virial coefficient  $\delta$  for the binary mixture:  $CO_2$ –air (see also SOP 24). He found that the temperature dependence of this parameter could be represented by the equation

$$
\frac{\delta (CO_2 - air)}{cm^3 mol^{-1}} = 57.7 - 0.118 \left(\frac{T}{K}\right)
$$
 (27)

where 273 < *T*/K < 313.

At 25°C (298.15 K),  $\delta$ (CO<sub>2</sub>-air) = 22.5 cm<sup>3</sup> mol<sup>-1</sup>.

#### **6. Salinity and the composition of sea water**

#### **6.1 The major ion composition of sea water**

Traditionally the parameter salinity, *S*, has been used to give a measure of the total dissolved solids in sea water. Originally this was intended to be an analytical parameter with the units  $g \text{ kg}^{-1}$ , but this direct approach was quickly superseded by the practice of estimating salinity from alternative, simpler procedures (*e.g.*, the measurement of chlorinity or conductivity ratio). The concept of a "constant composition of sea water" is then invoked to relate these parameters to salinity (UNESCO, 1966).

Although the idea of "constant composition of sea water" is recognized as being not strictly accurate, it is a useful device in simplifying the study of the various physicochemical properties of sea water. One example of this is the treatment of sea water as a two-component system—water and sea-salt—to represent its thermodynamic properties (*e.g.*, osmotic coefficient or density—see sections 3.2 and 4.2).

Another is the treatment of sea water as a constant ionic medium in which the thermodynamics of various chemical processes involving minor constituents can be studied (*e.g.*, gas solubility or acid–base equilibria).

Analytical results for the other major components of sea water are usually expressed relative to chlorinity (Table 2) and a standard mean chemical composition of sea water (Table 3) can be calculated using the equilibrium constants for the dissociation of water, carbonic and boric acids.

Component	<b>Relative</b> concentration <sup>a</sup>	mol $kg$ -soln <sup>-1</sup>	Reference	
Chloride	0.99889	0.54586	calculated from chlorinity	
Sulfate	0.1400	0.02824	Morris and Riley (1966)	
<b>Bromide</b>	0.003473	0.00084	Morris and Riley (1966)	
Fluoride	0.000067	0.00007	Riley (1965)	
Sodium	0.55661	0.46906	from charge balance	
Magnesium	0.06626	0.05282	Carpenter and Manella (1973)	
Calcium	0.02127	0.01028	Riley and Tongudai (1967)	
Potassium	0.0206	0.01021	Riley and Tongudai (1967)	
Strontium	0.00041	0.00009	Riley and Tongudai (1967)	
<b>Boron</b>	0.000232	0.000416	Uppström $(1974)$	
Total alkalinity		0.002400	average surface water	
$pH = 8.1$		$10^{-8.1}$	average surface water	

**Table 2** Standard mean analytical composition of sea water with *S* = 35 and chlorinity  $= 19.374$ .

<sup>a</sup> Expressed relative to the chlorinity (=  $S/1.80655$ ). Thus the total sulfate (molar mass 96.062 g) at a salinity, *S*, is given by

$$
S_{\rm T} = \left(\frac{0.1400}{96.062}\right) \times \left(\frac{S}{1.80655}\right) \text{mol kg-soln}^{-1}.
$$

<b>Species</b>	mol $kg$ -soln $^{-1}$	$g\ kg$ -soln <sup>-1</sup>	mol kg- $H_2O^{-1}$	g kg-H <sub>2</sub> $O^{-1}$
$Cl^{-}$	0.54586	19.3524	0.56576	20.0579
$SO_4^{2-}$	0.02824	2.7123	0.02927	2.8117
$Br^-$	0.00084	0.0673	0.00087	0.0695
$F^{-}$	0.00007	0.0013	0.00007	0.0013
$\mathrm{Na}^+$	0.46906	10.7837	0.48616	11.1768
$Mg^{2+}$	0.05282	1.2837	0.05475	1.3307
$Ca^{2+}$	0.01028	0.4121	0.01065	0.4268
$K^{\scriptscriptstyle +}$	0.01021	0.3991	0.01058	0.4137
$\text{Sr}^{2+}$	0.00009	0.0079	0.00009	0.0079
B(OH)	0.00032	0.0198	0.00033	0.0204
$B(OH)_{4}^{-}$	0.00010	0.0079	0.00010	0.0079
$CO2*$	0.00001	0.0004	0.00001	0.0004
HCO <sub>3</sub>	0.00177	0.1080	0.00183	0.1117
CO <sub>3</sub> <sup>2</sup>	0.00026	0.0156	0.00027	0.0162
$OH-$	0.00001	0.0002	0.00001	0.0002
sum of column	1.11994	35.1717	1.16075	36.4531
ionic strength	0.69734		0.72275	

**Table 3** Standard mean chemical composition of sea water  $(S = 35)$ .

#### **6.2 A simplified synthetic sea water recipe**

A simplified synthetic sea water recipe can be obtained from the recipe in Table 2 as follows: replace bromide, fluoride, and total alkalinity with chloride and replace strontium with calcium. The resultant composition (Table 4) is the basis of the synthetic sea water that has been used to determine a variety of equilibrium constants for use in sea water (*e.g.*, Dickson, 1990; Roy *et al.*, 1993).

<b>Species</b>	mol $kg$ -soln <sup>-1</sup>	$g\ kg$ -soln <sup>-1</sup>	mol kg- $H_2O^{-1}$	g kg-H <sub>2</sub> $O^{-1}$
$Cl^{-}$	0.54922	19.4715	0.56918	20.1791
$\mathrm{SO}_4^{2-}$	0.02824	2.7128	0.02927	2.8117
$Na+$	0.46911	10.7848	0.48616	11.1768
$Mg^{2+}$	0.05283	1.2840	0.05475	1.3307
$Ca^{2+}$	0.01036	0.4152	0.01074	0.4304
$K^+$	0.01021	0.3992	0.01058	0.4137
sum of column	1.11997	35.0675	1.16068	36.3424
ionic strength	0.69713		0.72248	

**Table 4** Simplified synthetic sea water composition  $(S = 35)^a$ .

a The composition here is very slightly different from that used by Dickson (1990) or Roy *et al.* (1993). This is the result of minor changes in the molar masses used to compute the various concentrations.

### **7. Equilibrium constants**

All the equations for the equilibrium constants presented here use concentrations expressed in moles per kilogram of solution. In addition, with the exception of that for bisulfate ion, all acid dissociation constants are expressed in terms of "total" hydrogen ion concentration (see Chapter 2).

#### **7.1 Solubility of carbon dioxide in sea water**

The equilibrium constant for the process

$$
CO2(g) \rightleftharpoons CO2*(aq),
$$
 (28)

*i.e.*,

$$
K_0 = [CO_2^*]/f(CO_2), \tag{29}
$$

is given by the expression (Weiss, 1974)

$$
\ln(K/k^{\circ}) = 93.4517 \left(\frac{100}{T/K}\right) - 60.2409 + 23.3585 \ln\left(\frac{T/K}{100}\right) + S \left[0.023517 - 0.023656\left(\frac{T/K}{100}\right) + 0.0047036\left(\frac{T/K}{100}\right)^{2}\right].
$$
 (30)

The fugacity of  $CO_2$  gas (see Chapter 2) is expressed in atm;  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At 
$$
S = 35
$$
 and  $t = 25^{\circ}$ C (298.15 K), ln  $(K_0/k^{\circ}) = -3.5617$ .

#### **7.2 Acid–base reactions in sea water**

#### *7.2.1 Bisulfate ion*

The equilibrium constant for the reaction

$$
HSO_4^- \rightleftharpoons H^+ + SO_4^-, \tag{31}
$$

*i.e.*,

$$
K_{\rm S} = \left[\rm H^+ \right]_F \left[\rm SO_4^{2-} \right] / \left[\rm HSO_4^- \right],\tag{32}
$$

is given by the expression (Dickson, 1990a)

$$
\ln(K_{\rm s}/k^{\circ}) = \frac{-4276.1}{(T/K)} + 141.328 - 23.093 \ln(T/K)
$$
  
+ 
$$
\left(\frac{-13856}{(T/K)} + 324.57 - 47.986 \ln(T/K)\right) \times \left(\frac{I}{m^{\circ}}\right)^{1/2}
$$
  
+ 
$$
\left(\frac{35474}{(T/K)} - 771.54 + 114.723 \ln(T/K)\right) \times \left(\frac{I}{m^{\circ}}\right)
$$
  
- 
$$
\frac{2698}{(T/K)} \left(\frac{I}{m^{\circ}}\right)^{3/2} + \frac{1776}{(T/K)} \left(\frac{I}{m^{\circ}}\right)^{2} + \ln(1 - 0.001005S)
$$
(33)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup> and hydrogen ion concentration is expressed on the "free" scale. The ionic strength is calculated from the expression

$$
I/m^{\circ} = \frac{19.924S}{1000 - 1.005S} \,. \tag{34}
$$

The term  $\ln(1 - 0.001005S)$  converts the value of  $K_S$  from mol kg-H<sub>2</sub>O<sup>-1</sup> (used in Dickson, 1990a) to mol kg-soln<sup>-1</sup>.

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15K), ln  $(K_S/k^{\circ}) = -2.30$ .

#### *7.2.2 Boric acid*

The equilibrium constant for the reaction

$$
B(OH)_3 + H_2O \rightleftharpoons H^+ + B(OH)_4^-, \tag{35}
$$

*i.e.*,

$$
K_{\rm B} = [H^+][B(OH)_4^-]/[B(OH)_3]
$$
\n(36)

is given by the expression (Dickson, 1990b)

$$
\ln\left(\frac{K_{\rm B}}{k^{\circ}}\right) = \frac{-8966.90 - 2890.53S^{1/2} - 77.942S + 1.728S^{3/2} - 0.0996S^{2}}{(T/K)} + (148.0248 + 137.1942S^{1/2} + 1.62142S) + (-24.4344 - 25.085S^{1/2} - 0.2474S)\ln(T/K) + 0.053105S^{1/2}(T/K) \tag{37}
$$

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>. This equation is in excellent agreement with the measurements made by Roy *et al.* (1993a); it is also in reasonable agreement with the results of Hansson (1973a).

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15 K),  $\ln(K_B/k^{\circ}) = -19.7964$ .

#### *7.2.3 Carbonic acid*

The equilibrium constants used here are based on the measurements of Mehrbach *et al*. (1973) and have been converted to the total hydrogen ion pH scale by Lueker *et al.* (2000) who also showed them to be in good agreement with direct measurements of  $p(CO_2)$ ,  $C_T$ , and  $A_T$ . These constants are in reasonable agreement with those measured by Roy *et al.* (1993b), Goyet and Poisson (1989), and Hansson (1973b).

The equilibrium constant for the reaction

$$
CO_2^*(aq) + H_2O(l) \rightleftharpoons H^*(aq) + HCO_3^-(aq), \qquad (38)
$$

*i.e.*,

$$
K_1 = [H^+][HCO_3^-]/[CO_2^*]
$$
 (39)

is given by the expression (Lueker *et al.*, 2000)

$$
\log_{10}(K_1/k^{\circ}) = \frac{-3633.86}{(T/K)} + 61.2172 - 9.67770 \ln(T/K)
$$
  
+0.011555S - 0.0001152S<sup>2</sup> (40)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At 
$$
S = 35
$$
 and  $t = 25^{\circ}$ C (298.15 K),  $log_{10}(K_1/k^{\circ}) = -5.8472$ .

The equilibrium constant for the reaction

$$
HCO3-(aq) \rightleftharpoons H+(aq) + CO32-(aq),
$$
\n(41)

*i.e.*,

$$
K_2 = [H^+][CO_3^{2-}]/[HCO_3^-],\tag{42}
$$

is given by the expression (Lueker *et al.,* 2000)

$$
\log_{10}(K_2/k^{\circ}) = \frac{-471.78}{(T/K)} - 25.9290 + 3.16967 \ln(T/K) + 0.01781S - 0.0001122S^2
$$
 (43)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15 K),  $log_{10}(K_2/k^{\circ}) = -8.9660$ .

#### *7.2.4 Hydrogen fluoride*

The equilibrium constant for the reaction

$$
HF(aq) \rightleftharpoons H^+(aq) + F^-(aq), \tag{44}
$$

*i.e.*,

$$
K_{\rm F} = [H^+][F^-]/[HF],\tag{45}
$$

is given by the expression (Perez and Fraga, 1987)

$$
\ln(K_{\rm F}/k^{\circ}) = \frac{874}{(T/K)} - 9.68 + 0.111S^{1/2}
$$
 (46)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>. This equation gives values that are in reasonable agreement with those suggested by Dickson and Riley (1979a).

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15K),  $\ln(K_F/k^{\circ}) = -6.09$ .

#### *7.2.5 Phosphoric acid*

The expressions below are from Millero  $(1995)^5$  and are a composite of measurements by Kester and Pytkowicz (1967), Dickson and Riley (1979b), and Johansson and Wedborg (1979).

 $\overline{a}$ 

 $5$  0.015 has been subtracted from the constant term in each of these expressions: (49), (52), (55) to convert—approximately—from the SWS pH scale (including HF) used by Millero (1995) to the "total" hydrogen ion scale used here.

The equilibrium constant for the reaction

$$
H_3PO_4(aq) \rightleftharpoons H^+(aq) + H_2PO_4^-(aq), \qquad (47)
$$

*i.e.*,

$$
K_{1P} = [H^+][H_2PO_4^-]/[H_3PO_4],
$$
\n(48)

is given by the expression

$$
\ln(K_{1P}/k^{\circ}) = \frac{-4576.752}{T/K} + 115.525 - 18.453 \ln(T/K)
$$

$$
+ \left(\frac{-106.736}{(T/K)} + 0.69171\right) S^{1/2} + \left(-\frac{0.65643}{(T/K)} - 0.01844\right) S \tag{49}
$$

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At *S* = 35 and *t* = 25°C (298.15 K),  $\ln(K_{1P}/k^{\circ})$  = -3.71.

The equilibrium constant for the reaction

$$
H_2PO_4^-(aq) \rightleftharpoons H^+(aq) + HPO_4^{2-}(aq), \qquad (50)
$$

*i.e.*,

$$
K_{2P} = [H^+][HPO_4^{2-}]/[H_2PO_4^-],
$$
\n(51)

is given by the expression

$$
\ln(K_{2P}/k^{\circ}) = \frac{-8814.715}{T/K} + 172.0883 - 27.927 \ln(T/K)
$$
  
+ 
$$
\left(\frac{-160.340}{(T/K)} + 1.3566\right) S^{1/2} + \left(\frac{0.37335}{(T/K)} - 0.05778\right) S
$$
 (52)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At *S* = 35 and *t* = 25°C (298.15 K),  $\ln(K_{2P}/k^{\circ})$  = -13.727.

The equilibrium constant for the reaction

$$
HPO42-(aq) \rightleftharpoons H+(aq) + PO43-(aq), \qquad (53)
$$

*i.e.*,

$$
K_{3P} = [H^+][PO_4^{3-}]/[HPO_4^{2-}],
$$
\n(54)

is given by the expression

$$
\ln(K_{3P}/k^{\circ}) = \frac{-3070.75}{(T/K)} - 18.141 + \left(\frac{17.27039}{(T/K)} + 2.81197\right) S^{1/2} + \left(\frac{-44.99486}{(T/K)} - 0.09984\right) S
$$
 (55)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>.

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15 K),  $\ln(K_{3P}/k^{\circ}) = -20.24$ .

#### *7.2.6 Silicic acid*

The equilibrium constant for the reaction

$$
\text{Si(OH)}_{4}\text{(aq)} \rightleftharpoons \text{H}^{+}\text{(aq)} + \text{SiO(OH)}_{3}\text{(aq)},\tag{56}
$$

*i.e.*,

$$
K_{\rm Si} = [H^+][\rm SiO(OH)_3^-]/[Si(OH)_4],\tag{57}
$$

is given by the expression (Millero, 1995)

$$
\ln(K_{\text{Si}}/k^{\circ}) = \frac{-8904.2}{(T/K)} + 117.385 - 19.334 \ln(T/K)
$$
  
+ 
$$
\left(\frac{-458.79}{(T/K)} + 3.5913\right) (I/m^{\circ})^{1/2} + \left(\frac{188.74}{(T/K)} - 1.5998\right) (I/m^{\circ})
$$
  
+ 
$$
\left(\frac{-12.1652}{(T/K)} + 0.07871\right) (I/m^{\circ})^2 + \ln(1 - 0.001005S)
$$
(58)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>. This expression is based on the results of Sjöberg *et al.* (1981) and the review by Baes and Mesmer (1976); 0.015 has been subtracted from the constant (see Footnote 5).

The ionic strength is calculated from the expression

$$
I/m^{\circ} = \frac{19.924S}{1000 - 1.005S} \approx 0.02S
$$
 (59)

The term  $ln(1 - 0.001005S)$  converts the value of  $K_{Si}$  from mol kg-H<sub>2</sub>O<sup>-1</sup> to mol kg-soln<sup>-1</sup>;  $m^{\circ}$  is as in equation (11).

At  $S = 35$  and  $t = 25^{\circ}$ C (298.15 K),  $\ln(K_{S_i}/k^{\circ}) = -21.61$ .

#### *7.2.7 Water*

The equilibrium constant for the reaction

$$
H_2O(l) \rightleftharpoons H^+(aq) + OH^-(aq),\tag{60}
$$

*i.e.*,

$$
K_{\rm w} = \left[\rm H^+ \right] \left[\rm OH^- \right],\tag{61}
$$

is given by the expression (Millero, 1995)

$$
\ln\left(K_{\rm w}/(k^{\circ})^2\right) = \frac{-13847.26}{(T/K)} + 148.9652 - 23.6521\ln(T/K) + \left(\frac{118.67}{(T/K)} - 5.977 + 1.0495\ln(T/K)\right) S^{1/2} - 0.01615S
$$
 (62)

where  $k^{\circ} = 1$  mol kg-soln<sup>-1</sup>. This expression is based on results published by Hansson (1973a), Culberson and Pytkowicz (1973), and Dickson and Riley (1979a). Note that 0.015 has been subtracted from the constant term (see Footnote 5).

At 
$$
S = 35
$$
 and  $t = 25^{\circ}$ C (298.15 K),  $ln(K_w/(k^{\circ})^2) = -30.434$ .

#### **7.3 Acid–base reactions in sodium chloride media**

The values given here are appropriate to a background medium with *C*(NaCl) = 0.7 mol kg-soln<sup>-1</sup> and at a temperature of  $25^{\circ}$ C.

#### *7.3.1 Carbonic acid*

The values given here are from Dyrssen and Hansson (1973).

The equilibrium constant for the reaction

$$
CO2*(aq) + H2O(l) \rightleftharpoons H+(aq) + HCO3-(aq),
$$
\n(63)

*i.e.*,

$$
K_1 = [H^+][HCO_3^-]/[CO_2^*],\tag{64}
$$

at  $C(NaCl) = 0.7$  mol kg-soln<sup>-1</sup> and  $t = 25$ °C is

$$
\ln(K_1/k^{\circ}) = -13.82 \,. \tag{65}
$$

The equilibrium constant for the reaction

$$
HCO3-(aq) \rightleftharpoons H+(aq) + CO32-(aq), \qquad (66)
$$

*i.e.*,

$$
K_2 = [H^*][CO_3^{2-}]/[HCO_3^-],\tag{67}
$$

at  $C(NaCl) = 0.7$  mol kg-soln<sup>-1</sup> and  $t = 25$ °C is

$$
\ln(K_2/k^{\circ}) = -21.97. \tag{68}
$$

#### *7.3.2 2-amino-2-hydroxymethyl-1,3-propanediol ("TRIS")*

The equilibrium constant for the reaction

$$
H3NC(CH2OH)3+(aq) \rightleftharpoons H+(aq) + H2NC(CH2OH)3(aq),
$$
 (69)

*i e.*,

$$
K_{\text{tris}} = [H^*][H_2NC(CH_2OH)_3]/[H_3NC(CH_2OH)_3^+],\tag{70}
$$

at  $C(NaCl) = 0.7$  mol kg-soln<sup>-1</sup> and  $t = 25$ °C is

$$
\ln (K_{TRS}/k^{\circ}) = -18.90 \tag{71}
$$

(Millero *et al.*, 1987).

#### *7.3.3 Water*

The equilibrium constant for the reaction

$$
H_2O(l) \rightleftharpoons H^+(aq) + OH^-(aq),\tag{72}
$$

*i.e.*,

$$
K_{\rm w} = \left[\rm H^+ \right] \left[\rm OH^- \right],\tag{73}
$$

at  $C(NaCl) = 0.7$  mol kg-soln<sup>-1</sup> and  $t = 25$ °C is

$$
\ln\left(K_{\rm w}/(k^{\circ})^2\right) = -31.71\tag{74}
$$

(Dyrssen and Hansson, 1973).

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